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Supplementary Information for

Spin transition in magnesiowüstite and ultralow thermal conduction in ultralow velocity zones

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21 **Note S1. Synchrotron X-ray diffraction on the (Mg_{0.25},Fe_{0.75})O, Fp75, at high**
22 **pressures**

23 We performed synchrotron X-ray diffraction (XRD) measurements on the Fp75 at
24 high pressures to ~89 GPa and room temperature at the beamline 12.2.2 of the Advanced
25 Light Source, Lawrence Berkeley National Laboratory. We prepared two sets of samples
26 within 200 μm culet diamond-anvil cells, loaded with silicone oil and neon as the pressure
27 medium, respectively, and ruby and gold chips as pressure calibrants. Supplementary Fig.
28 S1 shows the representative XRD patterns and spectra at selected pressures. When loaded
29 with silicone oil as the medium (the same as our separate thermal conductivity
30 measurements at room temperature), we found that the Fp75 remained a cubic structure at
31 16.8 GPa, while the (111) peak started to split at 23.5 GPa, suggesting a cubic to
32 rhombohedral structural distortion. Fei *et al.*¹ observed a cubic to rhombohedral structural
33 distortion in (Mg_{0.42},Fe_{0.58})O, Fp58, at ~44 GPa and made a phase diagram predicting the
34 onset pressures for such distortion as a function of FeO content in (Mg,Fe)O. Our XRD
35 observation is in good agreement with that suggested by Fei *et al.*¹, where the occurrence
36 of structural distortion in Fp75 would be at ~20–30 GPa. Since we did not observe a
37 significant change in the thermal conductivity of Fp75, Λ_{Fp75} , around this pressure range
38 (Fig. 1 of the main text), the cubic to rhombohedral structural distortion seems to have
39 minor effects on its thermal conductivity.

40 Supplementary Fig. S2 summarizes our present results for the pressure dependence of
41 the unit-cell-volume of Fp75 along with relevant literature results with different FeO
42 contents. Measurements with silicone oil (red circles) and with neon (red diamonds) as the
43 medium consistently show an increase in compressibility at ~55 GPa, suggesting the onset
44 of spin transition of iron. The different onset pressure and the pressure range for the spin
45 transition in (Mg,Fe)O between the present and literature studies can be attributed to the
46 different FeO content, characteristics of the sample (e.g., single-crystalline or
47 polycrystalline powder), and the pressure medium used in each experiment, where
48 nonhydrostatic stress induces the spin transition at lower pressures^{1,2}. These XRD results
49 suggest that the “sluggish” variation of the Λ_{Fp75} at $P > \sim 53$ GPa in the Fig. 1 of the main
50 text is presumably associated with the spin transition of iron. Though we did not perform
51 high-pressure XRD on the Fp25, given its well-documented pressure range of spin

52 transition^{1,2}, the decrease in Λ_{Fp25} between ~ 43 and 54 GPa (black symbols in the Fig. 1 of
53 the main text) is expected to be resulted from the spin transition.

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55 **Note S2. Minor effect of the thermal pressure on the Λ_{Fp25} and Λ_{Fp75} at high**
56 **temperature measurements**

57 During the high temperature measurements, the pressure within the sample chamber
58 inevitably is varied slightly due to the heating-induced thermal pressure. We, however,
59 expect that such pressure variation does not significantly affect the temperature dependence
60 we obtained (Fig. 3 of the main text). This is justified by the relatively weak pressure
61 dependences of low-spin Λ_{Fp25} and Λ_{Fp75} within the pressure range of experiments, see
62 their values in the low-spin state in Fig. 2 of the main text. For instance, the low-spin Λ_{Fp25}
63 at 573 K and $P=62\text{--}68$ GPa (black symbols in Fig. 2 of the main text) minorly varies by
64 $<10\%$, which is less than the estimated data uncertainty. Therefore, the pressure variation
65 ($62\text{--}68$ GPa) during the high temperature measurements (red symbols in Fig. 3a of the
66 main text) can be treated as an essentially constant pressure. This treatment also applies to
67 the low-spin Λ_{Fp75} , as the Λ_{Fp75} has an even weaker pressure dependence, see red symbols
68 at $P > \sim 60$ GPa in Fig. 2 of the main text.

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70 **Note S3. Electrically insulating Fp25 and Fp75 at deep mantle conditions**

71 The electrical conductivity σ of FeO wüstite and magnesiowüstite with FeO content
72 higher than $\sim 80\%$ may increase as the P - T conditions approach those at the Earth's core-
73 mantle boundary (CMB). For instance, the σ of $(\text{Mg}_{0.2}, \text{Fe}_{0.8})\text{O}$, Fp80, at CMB's high P - T
74 conditions was reported to increase to $\sim 10^{4.5}$ S m^{-1} , yet treated as an electrical insulator³.
75 The σ of FeO at the CMB condition becomes $\sim 10^5$ S m^{-1} , approaching that of an electrical
76 conductor⁴. However, since the σ of $(\text{Mg}, \text{Fe})\text{O}$ is proportional to its FeO content^{3,4}, our
77 samples (Fp75 and Fp25) with FeO content lesser than the Fp80 are expected to be more
78 electrically insulating. In other words, the electrons have minor contributions to the Λ_{Fp25}
79 and Λ_{Fp75} we presented in the Figs. 2 and 3 of the main text as well as the modelled profiles
80 throughout the whole lower mantle in Fig. 4 of the main text.

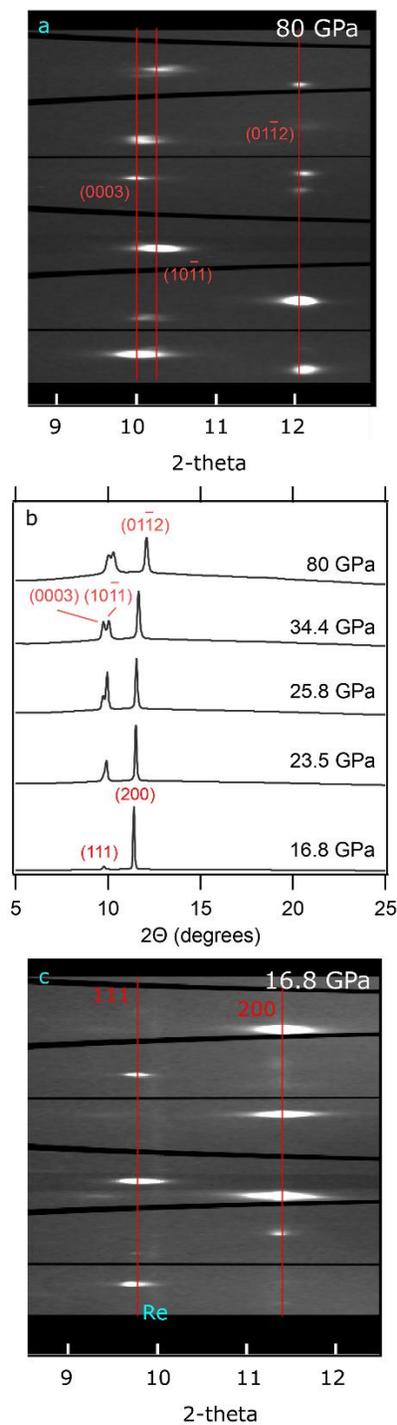
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82 **Note S4. Mössbauer spectroscopy for ferric irons in Fp25 and Fp75 at room**
83 **temperature**

84 The Mössbauer spectra of our Fp samples display a dominant doublet and a weak
85 shoulder on one side, which is very similar to those reported in the literature, e.g., Ref ⁵⁻⁷,
86 and therefore were fitted to one doublet and one singlet. All fitting parameters are treated
87 as free parameters. Although there is some variability, the fitted hyperfine parameters are
88 in good agreement with the literature data⁵⁻⁷. The $\text{Fe}^{3+}/\Sigma\text{Fe}$ ratio were calculated from the
89 relative areas, with the area of Fe^{3+} absorption constrained by the asymmetry of the main
90 doublets, making it independent of the fitting model. The $\text{Fe}^{3+}/\Sigma\text{Fe}$ in Fp25 is ~6% and in
91 Fp75 is ~12%, respectively (also see Supplementary Figure S5 and Table S2).

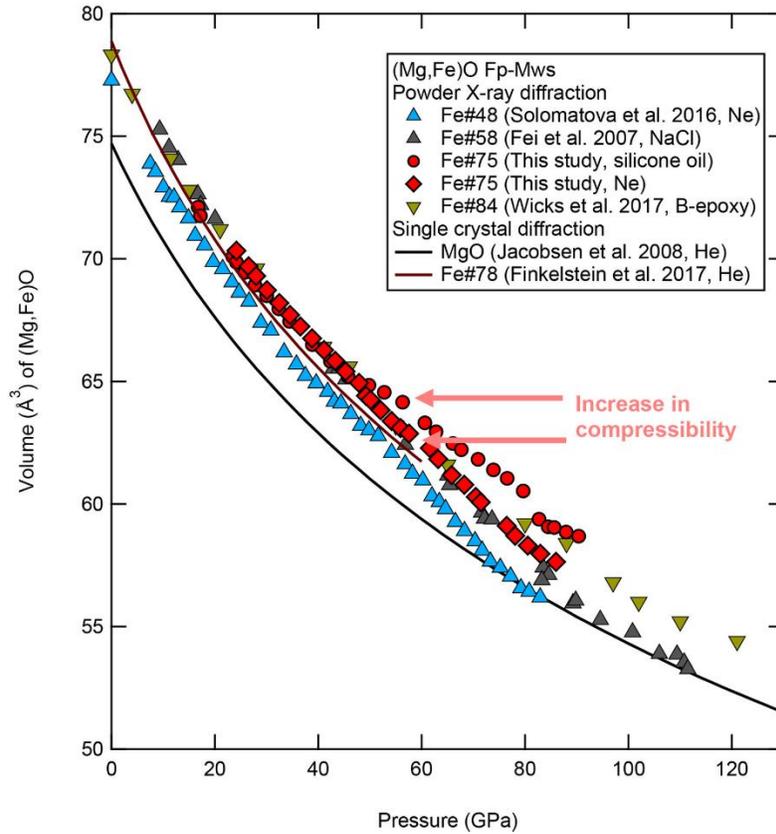
92 Mineral physics studies⁸ have indicated that a simple Fe_2O_3 could undergo a series of
93 phase transitions and reactions at deep mantle high P - T conditions, in which an oxygen-
94 rich fluid could be present in the deep lower mantle and locally oxidize surrounding
95 minerals. Since the (Mg,Fe)O is expected to be an important reservoir of ferric iron in the
96 lower mantle⁹, some amounts of ferric iron could be incorporated in Fe-rich
97 magnesiowüstite at deep lower mantle.

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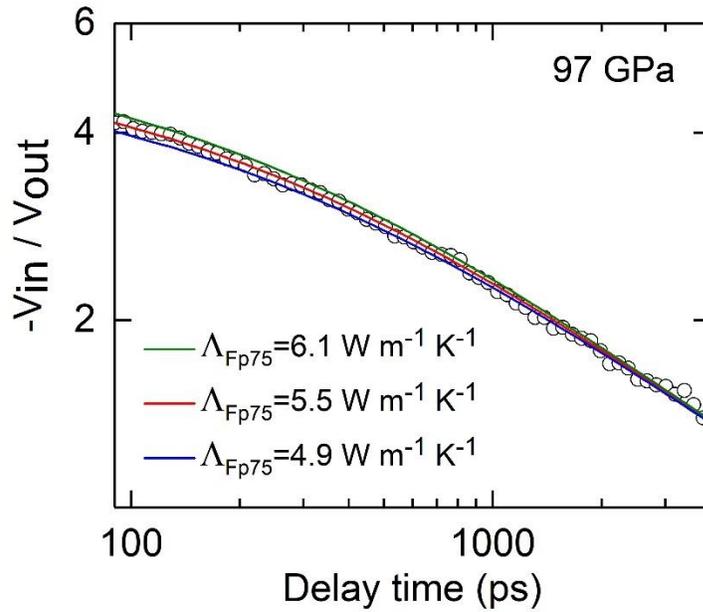
100 **Supplementary Figure S1.** Representative X-ray diffraction results for the Fp75
 101 compressed in silicone oil. (a) Unrolled 2D diffraction image obtained at 80 GPa indicating
 102 diffraction spots for rhombohedral structure with labeled Miller indices. (b) Integrated
 103 diffraction intensity as a function of diffraction angle 2θ for selected pressures. (c)
 104 Unrolled 2D diffraction image obtained at 16.8 GPa indicating diffraction spots for cubic
 105 B1 structure. Continuous background due to rhenium (Re) gasket is also observed at $\sim 10^\circ$.



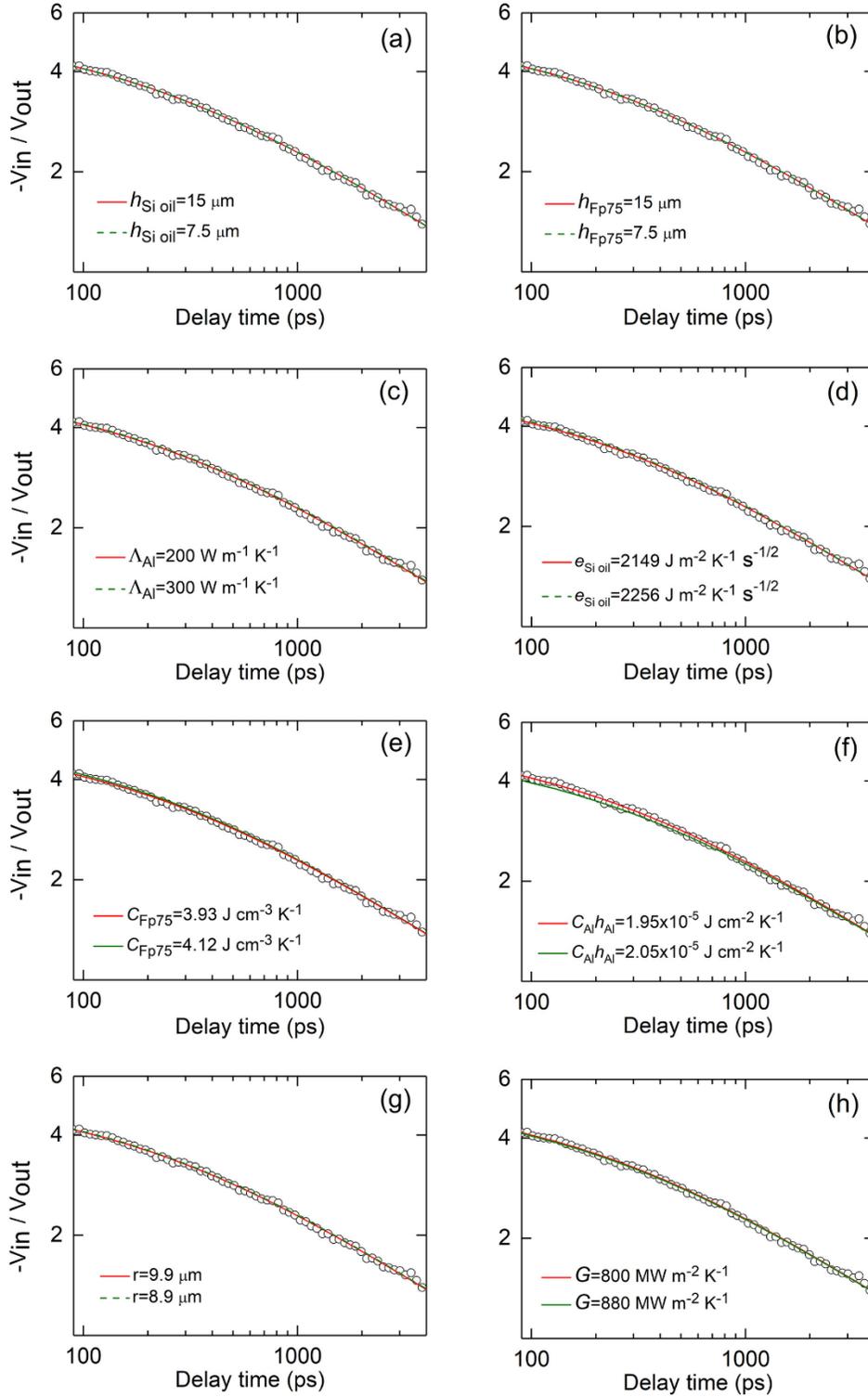
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108 **Supplementary Figure S2.** Unit cell volume for Fp75 (this study, red symbols) as a
 109 function of pressure compressed in silicone oil and Ne media computed from (200)
 110 diffraction peak. Observations from previous studies of less iron-rich ferropericlase-
 111 magnesiowüstite^{1,10} (up-pointing triangles) and more iron-rich compositions (down-
 112 pointing triangles) by powder¹¹ and single crystal^{12,13} diffraction (curves) are also shown
 113 for comparison. Our two sets of Fp75 data consistently show an increase in compressibility
 114 at ~55 GPa, suggesting the onset of spin transition.

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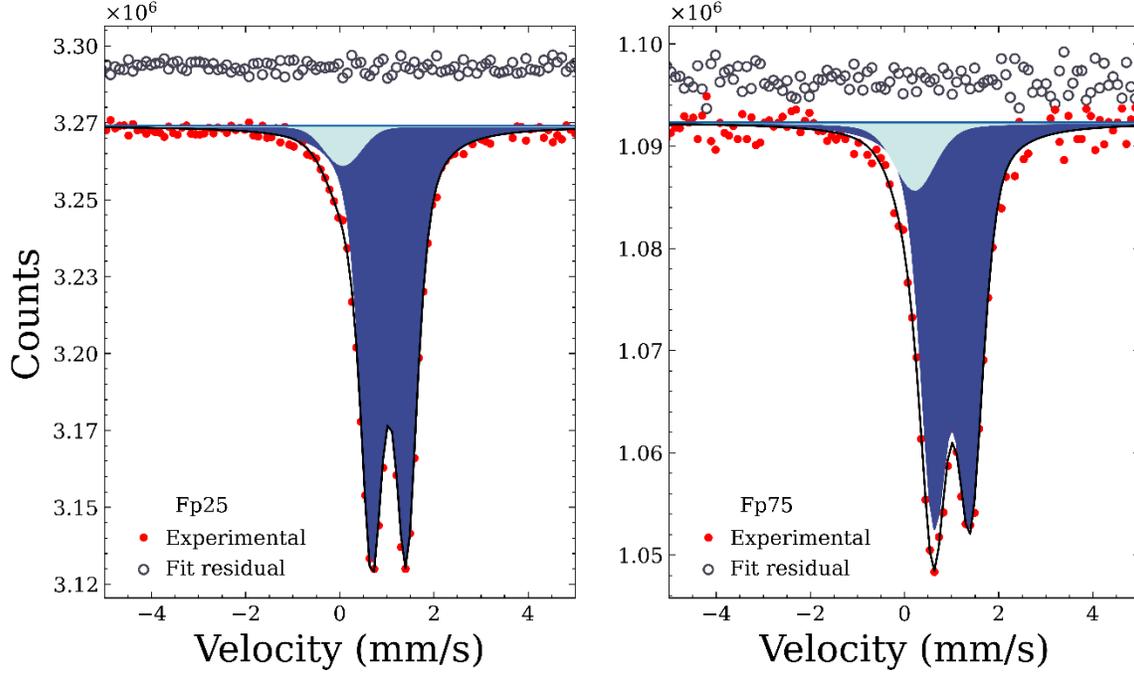
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 118 **Supplementary Figure S3.** Example TDTR data (open circles) for Fp75 at 97 GPa and
 119 room temperature along with thermal model calculations (color curves). The raw data are
 120 best-fitted with $\Lambda_{\text{Fp75}}=5.5 \text{ W m}^{-1} \text{ K}^{-1}$ (red curve) using input model parameters listed in
 121 Supplementary Table S1. The data ratio of $-V_{in}/V_{out}$ is most sensitive to the thermal
 122 conductivity of Fp75 at few hundred picosecond (ps) delay time^{14,15}. Minor changes by
 123 $\pm 10\%$ in the Λ_{Fp75} lead to clear mis-fits to the data, which indicate the high sensitivity of
 124 the thermal model fitting to our high-quality data with precise Λ_{Fep75} .
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Supplementary Figure S4. Standard sensitivity tests of the thermal model calculations to key parameters for Fp75 at 97 GPa and room temperature. Here the Fp75's thermal conductivity, Λ_{Fp75} , is fixed at $5.5 \text{ W m}^{-1} \text{ K}^{-1}$ as derived from Supplementary Fig. S3. (a) and (b) Even though the thickness of silicone oil ($h_{\text{Si oil}}$) and Fp75 sample (h_{Fp75}),

131 respectively, changes by 50%, the thermal model calculation remains the same, verifying
132 that their uncertainty does not affect the derived Λ_{Fp75} . (c) Since the Al's thermal
133 conductivity is fairly high at high pressures, an uncertainty as large as 50% does not
134 influence the Λ_{Fp75} . (d) If the thermal effusivity of the silicone oil pressure medium,
135 $e=(\Lambda_{\text{Si}}C_{\text{Si}})^{1/2}$, is off by 5%, the data can be re-fitted with a smaller Λ_{Fp75} of $5.4 \text{ W m}^{-1} \text{ K}^{-1}$,
136 i.e., translating $\sim 1.8\%$ error for the derived Λ_{Fp75} . (e) If the volumetric heat capacity of
137 Fp75, C_{Fp75} , is off by 5%, the data can be re-fitted also with a smaller Λ_{Fp75} of 5.4 W m^{-1}
138 K^{-1} , again, minorly translating $\sim 1.8\%$ error. (f) The uncertainty in the heat capacity of Al
139 per unit area (product of Al's volumetric heat capacity and thickness, $C_{\text{Al}}h_{\text{Al}}$)¹⁴ is the major
140 source of error for our data analysis. If it is off by 5%, for example, the data can be re-fitted
141 with a larger Λ_{Fp75} of $6.1 \text{ W m}^{-1} \text{ K}^{-1}$ (i.e., $\sim 10\%$ error). Finally, an example 10% error for
142 the (g) laser spot size and (h) thermal conductance of Al/silicone oil and Al/Fp75
143 interfaces, G , respectively, does not influence the thermal model calculations, indicating
144 that these uncertainties have essentially no effect on the Λ_{Fp75} .
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147 **Supplementary Figure S5.** Mössbauer spectra of Fp samples. The doublets are Fe²⁺ in
 148 Fp in dark blue, while Fe³⁺ in Fp is in cyan.

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152 **Supplementary Table S1.** Input parameters in the thermal model for Fp75 at 97 GPa and

153 room temperature in TDTR measurements

P (GPa)	C_{Fp75} (J cm ⁻³ K ⁻¹)	C_{Al} (J cm ⁻³ K ⁻¹)	h_{Al} (nm)*	$e=(\Lambda_{\text{Si}}C_{\text{Si}})^{1/2}$ (J m ⁻² K ⁻¹ s ^{-1/2})	r (μm)	$h_{\text{Si oil/Fp75}}$ (μm)	Λ_{Al} (W m ⁻¹ K ⁻¹)	G (MW m ⁻² K ⁻¹)
97	3.93	2.684	72.6	2149	9.9	15/15	200	800

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*In this experimental run, the Al's thickness at ambient pressure is 87.3 nm.

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C_{Fp75} : Fp75's heat capacity, C_{Al} : Al's heat capacity, h_{Al} : Al's thickness, e : silicone oil's

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thermal effusivity, r : laser spot size, $h_{\text{Si oil}}$: silicone oil's thickness, h_{Fp75} : Fp75's thickness,

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Λ_{Al} : Al's thermal conductivity, G : thermal conductance of Al/Fp75 and Al/silicone oil

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interfaces.

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161 **Supplementary Table S2.** Best fit hyperfine parameters of the Mössbauer spectra and the
 162 $\text{Fe}^{3+}/\Sigma\text{Fe}$ ratios in Fp25 and Fp75.

	Fe^{2+}			Fe^{3+}			$\text{Fe}^{3+}/\Sigma\text{Fe}$
	CS (mm/s)	QS (mm/s)	Area	CS (mm/s)	QS (mm/s)	Area	
Fp25	1.046(6)	0.729(8)	94(3)	0.06(9)	Not applicable	6(3)	6(3) %
Fp75	1.01(2)	0.77(3)	88(7)	0.2(1)	Not applicable	12(7)	12(7) %

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Supplementary References

- 167 1. Fei, Y. *et al.* Spin transition and equations of state of (Mg,Fe)O solid solutions.
 168 *Geophys Res Lett* **34**, L17307 (2007).
 169 2. Lin, J.-F., Speziale, S., Mao, Z. & Marquardt, H. Effects of the Electronic Spin
 170 Transitions of Iron in Lower Mantle Minerals: Implications for Deep Mantle
 171 Geophysics and Geochemistry. *Reviews of Geophysics* **51**, 244–275 (2013).
 172 3. Ohta, K. *et al.* Highly conductive iron-rich (Mg,Fe)O magnesiowüstite and its
 173 stability in the Earth’s lower mantle. *J Geophys Res Solid Earth* **119**, 4656–4665
 174 (2014).
 175 4. Ho, W. G. D. *et al.* Quantum critical phase of FeO spans conditions of Earth’s
 176 lower mantle. *Nat Commun* **15**, 3461 (2024).
 177 5. McCammon, C. A., Stachel, T. & Harris, J. W. Iron oxidation state in lower
 178 mantle mineral assemblages II. Inclusions in diamonds from Kankan, Guinea.
 179 *Earth Planet Sci Lett* **222**, 423–434 (2004).
 180 6. Dobson, D. P., Cohen, N. S., Pankhurst, Q. A. & Brodholt, J. P. A convenient
 181 method for measuring ferric iron in magnesiowüstite (MgO-Fe_{1-x}O). *American*
 182 *Mineralogist* **83**, 794–798 (1998).
 183 7. Otsuka, K., Longo, M., McCammon, C. A. & Karato, S. ichiro. Ferric iron content
 184 of ferroperricite as a function of composition, oxygen fugacity, temperature and
 185 pressure: Implications for redox conditions during diamond formation in the lower
 186 mantle. *Earth Planet Sci Lett* **365**, 7–16 (2013).
 187 8. Bykova, E. *et al.* Structural complexity of simple Fe₂O₃ at high pressures and
 188 temperatures. *Nat Commun* **7**, 10661 (2016).
 189 9. Stixrude, L. & Lithgow-Bertelloni, C. Thermodynamics of mantle minerals – III:
 190 the role of iron. *Geophys J Int* **237**, 1699–1733 (2024).
 191 10. Solomatova, N. V. *et al.* Equation of state and spin crossover of (Mg,Fe)O at high
 192 pressure, with implications for explaining topographic relief at the core-mantle
 193 boundary. *American Mineralogist* **101**, 1084–1093 (2016).
 194 11. Wicks, J. K., Jackson, J. M., Sturhahn, W. & Zhang, D. Sound velocity and density
 195 of magnesiowüstites: Implications for ultralow-velocity zone topography. *Geophys*
 196 *Res Lett* **44**, 2148–2158 (2017).

- 197 12. Finkelstein, G. J. *et al.* Single-crystal equations of state of magnesiowüstite at high
198 pressures. *American Mineralogist* **102**, 1709–1717 (2017).
- 199 13. Jacobsen, S. D. *et al.* Compression of single-crystal magnesium oxide to 118 GPa
200 and a ruby pressure gauge for helium pressure media. *American Mineralogist* **93**,
201 1823–1828 (2008).
- 202 14. Zheng, X., Cahill, D. G., Krasnochtchekov, P., Averbach, R. S. & Zhao, J. C.
203 High-throughput thermal conductivity measurements of nickel solid solutions and
204 the applicability of the Wiedemann-Franz law. *Acta Mater* **55**, 5177–5185 (2007).
- 205 15. Cahill, D. G. & Watanabe, F. Thermal conductivity of isotopically pure and Ge-
206 doped Si epitaxial layers from 300 to 550 K. *Phys. Rev. B* **70**, 235322 (2004).
- 207